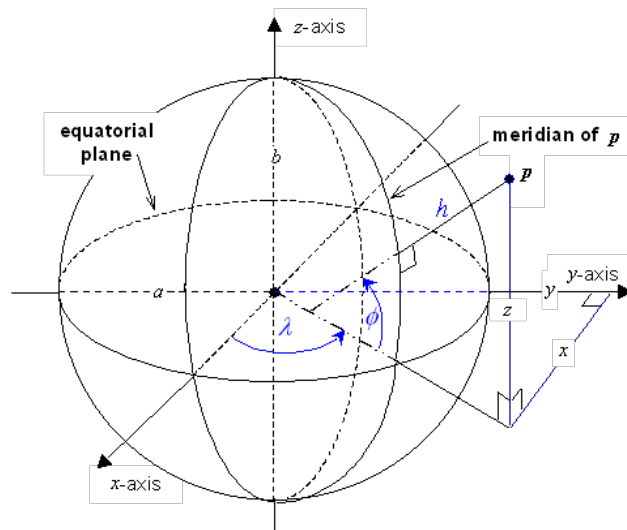


$X Y Z \leftrightarrow \varphi \lambda h$

di Michele T. Mazzucato

L'origine del sistema di coordinate cartesiane ortogonali nello spazio X, Y, Z coincide con il centro geometrico dell'ellissoide utilizzato come superficie di riferimento per le coordinate di geografiche ellissoidiche φ , λ , h . L'asse Z coincide con l'asse minore o di rotazione dell'ellissoide ed è positivo verso il Polo Nord, l'asse X giace nel piano equatoriale XY (origine delle latitudini) nell'intersezione con il piano meridiano XZ di Greenwich (origine delle longitudini) ed è positivo verso il meridiano di Greenwich e l'asse Y giace nel piano equatoriale XY, ortogonale all'asse X ed è destrorso positivo.



da <https://standards.sedris.org/18026/index.htm>

Il *passaggio diretto* dalle coordinate geografiche ellissoidiche φ , λ , h a quelle cartesiane ortogonali nello spazio X, Y, Z non presenta particolari difficoltà. Esso avviene mediante le seguenti relazioni:

$$X = \left(\frac{a}{\sqrt{1 - e^2 \sin^2 \varphi}} + h \right) \cos \varphi \cos \lambda \quad Y = \left(\frac{a}{\sqrt{1 - e^2 \sin^2 \varphi}} + h \right) \cos \varphi \sin \lambda$$

$$Z = \left[\frac{a}{\sqrt{1 - e^2 \sin^2 \varphi}} (1 - e^2) + h \right] \sin \varphi$$

in cui a è il semiasse maggiore ed e l'eccentricità prima numerica dell'ellissoide di riferimento geodetico

Diversamente accade per il *passaggio inverso* dalle coordinate cartesiane ortogonali nello spazio X, Y, Z a quelle geografiche ellissoidiche φ , λ , h .

Mentre, per la determinazione della longitudine λ si utilizza, generalmente, la relazione

$$\lambda = \arctan \frac{Y}{X} \pm 180^\circ$$

oppure la formula di Krakiwsky & Vanicek (1982) che la determina senza bisogno di definirne il segno e quindi il suo fuso, numericamente più stabile in vicinanza dei poli

$$\lambda = 2 \arctan \frac{Y}{X + \sqrt{X^2 + Y^2}}$$

Krakiwsky & Vaniček (1982)

oppure la formula di Vermeille (2004) con una stabilità numerica maggiore rispetto alla precedente

$$\lambda = \frac{\pi}{2} - 2 \arctan \frac{X}{\sqrt{X^2 + Y^2} + Y} \text{ per } Y \geq 0$$

$$\lambda = \frac{\pi}{2} + 2 \arctan \frac{X}{\sqrt{X^2 + Y^2} - Y} \text{ per } Y < 0$$

Vermeille (2004)

e per la determinazione dell'altitudine h , una volta ottenuta la latitudine φ , si utilizza

$$h = \frac{\sqrt{X^2 + Y^2}}{\cos \varphi} - N$$

oppure altre formule alternative come

$$h = \frac{Z}{\sin \varphi} - N(1 - e^2)$$

Bencini (1968)

$$h = \sqrt{X^2 + Y^2} \cos \varphi + Z \sin \varphi - a \sqrt{(1 - e^2 \sin^2 \varphi)}$$

Bowring (1985)

la determinazione della latitudine φ , invece, risulta più complicata in quanto nessuna relazione semplice la collega a X , Y , Z . Essa richiede tecniche indirette (o iterative) o metodi diretti (formule chiuse) di gran lunga più complessi di quelle indirette. In letteratura numerosi sono gli studi e le procedure proposte da vari autori (vedere tabella 1).

Simple Iteration

procedimento descritto in vari testi di geodesia come in Bomford, Guy (1899-1996) *Geodesy* (4ª ed. 1980, p. 679)

$$\tan \varphi_1 = \frac{Z + \frac{a}{\sqrt{(1 - e^2 \sin^2 \varphi_0)}} e^2 \sin \varphi_0}{\sqrt{X^2 + Y^2}}$$

il procedimento iterativo inizia con il valore di φ fornito dalla $\tan \varphi_0 \approx \frac{Z(1 - e'^2)}{\sqrt{X^2 + Y^2}}$

in cui e' è l'eccentricità seconda numerica dell'ellissoide di riferimento geodetico

Tabella 1

**Conversione cartesiane geocentriche ↔ geografiche ellissoidiche.
Elenco alfabetico, non esaustivo, per autore**

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NOTE

- alcune forme dirette o chiuse (non iterative) mantengono tale proprietà solo per $h=0$ (punto giacente sulla superficie di riferimento ellissoidica);

- Vincenty (1985) afferma che la prima soluzione a questo problema è stata data da Dörrie, Heinrich (1873-1955) in *Kubische und biquadratische Gleichungen* (1948);

